https://doi.org/10.46861/bmp.29.108

New data on sulphosalts from hydrothermal siderite-type veins in the Spišsko-gemerské rudohorie Mts. (eastern Slovakia): 1. Nuffieldite and aikinite from Slovinky-Došťavná vein

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ŠTEVKO M, SEJKORA J, DOJČANSKÝ Ľ (2021) New data on sulphosalts from hydrothermal siderite-type veins in the Spišsko-gemerské rudohorie Mts. (eastern Slovakia): 1. Nuffieldite and aikinite from Slovinky-Došťavná vein. Bull Mineral Petrolog 29(1): 108-114 ISSN 2570-7337

Abstract

A new occurrence of nuffieldite was recently discovered in a siderite-type hydrothermal vein with sulphides in Došťavná near Slovinky, Spišsko-gemerské rudohorie Mts., Spišská Nová Ves Co., Košice Region, Slovakia. It forms lead-gray acicular crystals reaching up to 1 cm in size or aggregates up to 2.5 cm, which are enclosed in quartz-siderite matrix together with chalcopyrite, pyrite, tourmaline and chlorite. Nuffieldite is frequently replaced by minor aikinite, galena and native bismuth. The refined unit-cell parameters of nuffieldite (for the orthorhombic space group *Pbnm*) are: a 14.5313(16) Å, b 21.454(2) Å, c 4.0500(6) Å and V 1262.58(19) Å³. The average (n=145 analyses) empirical formula of nuffieldite from Slovinky-Došťavná based on Pb+Bi+Sb = 5 *apfu* is corresponding to $Cu_{1.30}Pb_{2.00}Bi_{2.00}(Pb_{0.30}Bi_{0.30}Sb_{0.40})_{1.00}(S_{7.12}Se_{0.03})_{7.15}$. Aikinite forms microscopic, anhedral to subhedral grains or aggregates replacing nuffieldite. The average (n=29) empirical formula of studied aikinite based on (Cu+Pb)/2+(Sb+Bi) = 8 *apfu* is Pb_{3.83}Cu_{3.76}(Bi_{4.12} Sb_{0.09})_{4.21}(S_{1.240}Se_{0.03})_{1.205}.

Key words: nuffieldite, aikinite, sulphosalts, X-ray powder data, chemical composition, Došťavná vein, Slovinky, Spišsko-gemerské rudohorie Mts., Slovak Republic

Received 4. 2. 2021; accepted 27. 5. 2021

Introduction

The Spišsko-gemerské rudohorie Mts. represent one of the most important accumulations of ore deposits in whole Carpathian mountain range. There are more than 1200 hydrothermal ore veins known within this relatively small area, with two major types of mineralization: siderite -type carbonate-quartz veins with sulphides (extensively exploited in Dobšiná, Štítnik, Rákoš, Rožňava, Drnava, Rudňany, Novoveská Huta, Hnilčík, Henclová, Prakovce, Gelnica, Slovinky, Medzev etc.) and quartz-stibnite veins (Betliar, Čučma, Štofova dolina, Helcmanovce, Poproč or Zlatá Idka). Furthermore, Sn-Mo-W bearing greisens or hydrothermal veins, hydrothermal veins with U-REE mineralization as well as strata-bound VMS pyrite-Cu-Pb-Zn ore mineralization and hydrothermal-metasomatic bodies of siderite and magnesite±talc are present (Varček 1962; Chovan et al. 1994; Grecula et al. 1995; Rojkovič 1997).

All of the above mentioned types of ore mineralization contain various sulphosalts mostly as accessory ore minerals. Abundant presence of minerals of tetrahedritetennantite series (especially Fe, Zn and locally also Hg dominant members) is very typical feature of the siderite -type veins (e.g. Bernard 1958, 1961; Varček 1957, 1959, 1960; Novák 1959, 1967; Trdlička 1967; Háber 1980; Cambel et al. 1985; Peterec 1990; Miškovic 1991; Háber et al. 1993; Grecula et al. 1995; Antal 2002a, b; Pršek

2008; Pršek, Biroň 2007; Pršek, Lauko 2009; Števko et al. 2015; Mikuš et al. 2018; Števko, Sejkora 2020). Bi sulphosalts are also quite common, especially minerals of bismuthinite-aikinite series (e.g. Paděra et al. 1955; Kupčík et al. 1969; Hurný, Krištín 1978; Mumme, Žák 1983; Antal 1991; Beňka, Siman 1994; Pršek 2008; Števko et al. 2015; Mikuš et al. 2018, 2019) and kobellite homologous series (e.g. Trdlička, Kupka 1957; Hak, Kupka 1958; Novák 1961; Trdlička et al. 1962; Václav 1964; Zábranský, Radzo 1966; Háber, Streško 1969; Háber 1980; Jeleň 1991; Pršek 2008; Pršek, Peterec 2008; Mikuš et al. 2018, 2019). Other Bi sulphosalts like cosalite (Bernard 1964, Háber 1980), galenobismutite (Antal 1991, Pršek 2008), jaskólskiite (Pršek, Biroň 2007), nuffieldite (Pršek et al. 2006) or wittichenite (Háber 1978; Kozub et al. 2011) are rare. Chalcostibite is infrequent too (Sejkora et al. 2011; Mikuš et al. 2018). Unusual assemblage of Ag-Bi sulphosalts (matildite, gustavite and benjaminite) was recently described from Kobaltová vein near Medzev by Mikuš et al. (2019). The most common Pb sulphosalts at the siderite-type veins are bournonite, jamesonite and boulangerite (e.g. Zimányi 1914; Novák 1962; Trdlička 1967; Kupčík et al. 1969; Háber 1980; Miškovic 1990; Pršek, Biroň 2007; Pršek, Peterec 2008; Sejkora et al. 2011; Mikuš et al. 2018, 2019), whereas berthierite and garavellite (Mikuš et al. 2018), meneghinite (Beňka, Siman 1994) or zinkenite and scainiite (Sejkora et al. 2011) are scarce. Rare Hg sulphosalts, marruccite (Sejkora et al. 2011) and grumiplucite (Števko et al. 2015) were also recently identified.

In this series of papers, we would like to present new data on various sulphosalts from newly discovered localities in the Spišsko-gemerské rudohorie Mts.

Geological setting

The Došťavná vein is located in the Došťavná valley, around 7 km SW of the Slovinky village and 1 km NW of the Orlí vrch hill (1043 m a.s.l.) in the Spišsko-gemerské rudohorie Mts., Spišská Nová Ves Co., Košice Region, Slovakia. Samples of ore mineralization with sulphosalts were collected at the dump of abandoned adit (Fig. 1) located in the eastern part of the Došťavná vein/vein system. GPS coordinates (WGS84) of this dump are: 48.843134° N a 20.768606° E, 741 m asl.

The Došťavná vein is one of the minor siderite-type hydrothermal carbonate-quartz ore veins with sulphides, explored for Cu ores mainly in 19th century. According to Grecula et al. (1995) it is belonging to the Zlatá vein zone. Based on our recent detailed mapping of old mining works it is clear that rather than one vein, a system of se-

veral subparallel hydrothermal veins is developed in the Doštavná valley area. The E-W to NE-SW trending hydrothermal veins are mostly up to 0.6 m thick and are hosted by Early Paleozoic rocks of the Gelnica group, predominantly metarhyolite tuffs and phyllites of the Vlachovo Formation (Hurný 1973, 1974; Bajaník et al. 1984; Grecula et al. 1995). Mineralogy of the Došťavná vein was studied by Hurný (1973, 1974) and Hurný, Krištín (1978). The dominant gangue minerals are quartz, siderite and ankerite accompanied by minor amounts of chlorite, tourmaline, dickite and xenotime. Chalcopyrite and pyrite are the most frequent ore minerals associated with minor aikinite, hematite, magnetite, tetrahedrite and microscopic inclusions of native bismuth (Hurný 1973, 1974; Hurný, Krištín1978).

Analytical methods

The powder X-ray diffraction data of nuffieldite were collected on a Bruker D8 Advance diffractometer (National Museum, Prague, Czech Republic) with a solid-state 1D LynxEye detector using Cu*K* α radiation and operating at 40 kV and 40 mA. The powder patterns were collected using Bragg-Brentano geometry in the range 3 - 70° 2 θ , in 0.01° steps with a counting time of 20 s per step. Positions



Fig. 1 View on dump where samples with nuffieldite were collected. Photo by M. Števko, August 2020.





Fig. 3 Homogenous acicular crystals of nuffieldite (light gray) with galena (white) in quartz (black). BSE image by J. Sejkora.

Fig. 2 Aggregate of acicular crystals of nuffieldite enclosed with chalcopyrite in quartz. Nuffieldite aggregate is 2.5 cm long. Photo by P. Škácha.



Fig. 4 Aggregates of nuffieldite (gray) with galena (white) in quartz (black). BSE image by J. Sejkora.

and intensities of reflections were found and refined using the PearsonVII profile-shape function with the ZDS program package (Ondruš 1993) and the unit-cell parameters were refined by the leastsquares algorithm implemented by Burnham (1962). The experimental powder pattern was indexed in line with the calculated values of intensities obtained from the crystal structure of nuffieldite (Moëlo et al. 1997), based on Lazy Pulverix program (Yvon et al. 1977). The chemical analyses of nuffieldite and aikinite were performed using a Cameca SX100 electron microprobe (National Museum, Prague, Czech Republic) operating in the wave-dispersive (WDS) mode (25 kV, 20 nA and 0.7 μ m wide beam). The following standards and X-ray lines were used to minimize line overlaps: Ag (AgLa), Bi₂Se₃ (Bi*M* β), CdTe (CdLa), Co (CoKa), CuFeS₂ (CuKa, SKa), FeS₂ (FeKa), GaAs (GaLa), Ge (GeLa), HgTe (HgLa),



Fig. 5 Nuffieldite (Nf) replaced by aikinite (Ai), galena (Gn) and native bismuth (tiny white inclusions). BSE image by J. Sejkora.



Fig. 6 Nuffieldite (light gray) replaced by aikinite (dark gray) and native bismuth (white). BSE photo by J. Sejkora.

I _{obs.}	d _{obs.}	d _{calc.}	h	k	1	I _{obs.}	d _{obs.}	d _{calc.}	h	k	Ι	I _{obs.}	d _{obs.}	d _{calc.}	h	k	1
9.3	7.286	7.266	2	0	0	7.4	2.6373	2.6372	1	8	0	6.2	1.8204	1.8206	7	2	1
0.3	6.904	6.882	2	1	0	2.1	2.5903	2.5899	3	7	0	3.2	1.8101	1.8076	3	3	2
1.1	5.370	5.363	0	4	0	33.5	2.5488	2.5483	4	6	0	2.9	1.7968	1.7970	6	6	1
2.9	5.102	5.097	2	3	0	9.5	2.5299	2.5295	4	3	1	3.0	1.7915	1.7909	8	2	0
2.7	4.730	4.725	3	1	0	33	25154	2.5165	3	5	1	2.7	1.7890	1.7882	4	9	1
9.5	4.319	4.315	2	4	0	5.5	2.5154	2.5148	2	6	1	0.5	1.7707	1.7688	4	0	2
6.6	4.119	4.115	1	5	0	9.6	2.4080	2.4101	1	7	1	6.1	1.7615	1.7628	4	1	2
66.1	4.014	4.010	3	3	0	74	2 2620	2.3624	6	2	0	0.0	1.7615	1.7620	0	6	2
2.8	3.791	3.789	0	2	1	7.4	2.3029	2.3612	5	0	1	5.8	1.7465	1.7467	7	4	1
60.8	3.696	3.695	2	5	0	10.0	2.3432	2.3470	5	1	1	11	1 7197	1.7184	4	11	0
24.3	3.670	3.666	1	2	1	15.5	2.3064	2.3060	5	2	1	4.4	1.7 107	1.7187	7	7	0
45.8	3.635	3.633	4	0	0	1.5	2.2936	2.2939	6	3	0	2.4	1.7126	1.7130	3	5	2
100.0	2 500	3.582	4	1	0	2.8	2.2883	2.2878	4	5	1	3.7	1.7075	1.7080	2	11	1
100.0	3.560	3.576	0	6	0	9.1	2.2359	2.2360	0	8	1	3.4	1.6986	1.6989	6	9	0
10.5	3.491	3.490	2	1	1	7.8	2.2095	2.2100	1	8	1	2.5	1.6775	1.6775	5	9	1
7.3	3.443	3.441	4	2	0	5.6	2.1818	2.1819	3	7	1	4.4	1.6523	1.6524	8	1	1
14.1	3.428	3.425	1	3	1	5.9	2.1600	2.1611	5	4	1	6 9	1 6/11	1.6413	7	6	1
8.6	3.361	3.360	2	2	1	7.1	2.1379	2.1371	2	8	1	0.0	1.0411	1.6397	1	13	0
16.4	3.213	3.212	3	5	0	53	2 1000	2.1089	5	7	0	2.8	1.6348	1.6353	4	5	2
35.4	3.172	3.171	2	3	1	5.5	2.1090	2.1091	6	5	0	10	1 6054	1.6061	1	8	2
16.0	3.009	3.008	4	4	0	2.9	2.0572	2.0575	2	10	0	1.9	1.0054	1.6059	6	10	0
8.7	3.000	2.999	1	7	0	12.7	2.0394	2.0406	6	2	1	2.6	1.5848	1.5854	4	6	2
12.0	2.985	2.984	3	2	1	12.5	2.0247	2.0250	0	0	2	11	1 5450	1.5460	8	5	1
9.9	2.954	2.953	2	4	1	0	1 0029	1.9930	4	9	0	4.4	1.5459	1.5461	9	4	0
17.4	2.886	2.887	1	5	1	0.9	1.9950	1.9936	7	3	0	1 1	1 5221	1.5324	0	14	0
10.6	2.850	2.850	3	3	1	3.5	1.9764	1.9768	2	9	1	1.1	1.5521	1.5320	4	7	2
7.6	2.824	2.824	2	7	0	12.5	1.9708	1.9709	5	8	0	3.4	1.5193	1.5199	1	13	1
16.2	2.806	2.805	5	2	0	13.5	1.9363	1.9360	7	4	0	4.7	1.4957	1.4961	9	1	1
35.7	2.773	2.773	4	5	0	9.1	1.8909	1.8913	3	9	1	2.4	1.4911	1.4918	7	10	0
22.9	2.730	2.729	2	5	1	26.0	1.8833	1.8837	2	11	0	2.5	1.4572	1.4575	3	13	1
8.6	2.689	2.688	3	4	1	10.7	1.8703	1.8705	5	7	1						
10.6	2.681	2.682	0	8	0	6.1	1.8431	1.8431	5	9	0						

InAs (InLa), Mn (MnKa), NaCl (ClKa), NiAs (AsLB), Ni (NiKa), PbS (PbMa), PbSe (SeL β), PbTe (TeLa), Sb₂S₃ (SbLa), Sn (SnLa) and ZnS (ZnKa). Contents of the above-listed elements, which are not included in the tables, were analysed quantitatively, but their contents were below the detection limit (ca. 0.03 - 0.05 wt. % for individual elements). Raw intensities were converted to the concentrations of elements using automatic "PAP" matrix-correction software (Pouchou, Pichoir 1985). The hypothetical percentage of aikinite end member n_{aik} in aikinite was calculated according procedure proposed by Makovicky, Makovicky (1978).

Results and discussion

Nuffieldite is relatively common mineral at the studied locality. It forms lead-gray acicular crystals reaching up to 1 cm in size, which are enclosed together with chalcopyrite, pyrite, chlorite and tourmaline in quartz-siderite gangue. Aggregates of nuffieldite crystals (Fig. 2) reaching up to 2.5 cm have also been rarely observed. They are homogenous in BSE (Fig. 3, 4) and crystals or aggregates of nuffieldite are frequently replaced by minor amounts of aikinite, galena and native bismuth (Fig. 5, 6). The obtained experimental powder diffraction data set

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 Table 2 Unit-cell parameters of nuffieldite (for the orthorhombic space group Pbnm)

		<i>a</i> [Å]	b [Å]	c [Å]	V [ų]
Slovinky	this paper	14.5313(16)	21.454(2)	4.0500(6)	1262.58(19)
Les Houches	Moëlo et al. (1997)	14.4949(23)	21.4195(49)	4.0420(15)	1254.93(95)
Maleevskoe	Mozgova et al. (1994)	14.602	21.375	4.026	1256.6
Alice Arm	Kohatsu, Wuensch (1973)	14.387(7)	21.011(15)	4.046(6)	1223.0

 Table 3 Representative WDS analyses of nuffieldite from Slovinky (wt.%)

ky-Doštavná.

	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15
Pb	36.16	37.02	35.67	36.44	35.53	35.85	36.59	36.54	36.37	35.60	34.99	35.60	35.90	36.35	36.50
Cu	6.20	6.29	6.13	6.40	6.09	6.38	6.28	6.30	6.25	6.26	6.07	5.95	6.23	6.31	6.29
Sb	3.76	4.03	3.48	3.59	3.19	3.80	4.10	4.09	3.77	3.72	3.10	3.26	3.64	3.47	4.16
Bi	36.10	35.50	37.73	36.25	37.40	36.04	35.92	35.61	36.12	36.55	37.07	36.92	36.82	36.66	35.55
S	17.33	17.32	17.44	17.44	17.37	17.30	17.42	17.45	17.24	17.34	16.90	16.93	17.27	17.23	17.30
Se	0.14	0.18	0.20	0.00	0.19	0.26	0.10	0.13	0.00	0.10	0.57	0.55	0.23	0.15	0.11
total	99.69	100.34	100.65	100.12	99.77	99.63	100.41	100.12	99.75	99.57	98.70	99.21	100.09	100.17	99.91
Pb	2.308	2.341	2.258	2.321	2.276	2.297	2.311	2.318	2.314	2.277	2.272	2.289	2.284	2.312	2.315
Cu	1.290	1.297	1.265	1.329	1.272	1.333	1.293	1.303	1.296	1.306	1.285	1.248	1.292	1.309	1.301
Sb	0.408	0.434	0.375	0.389	0.348	0.414	0.441	0.442	0.408	0.405	0.342	0.357	0.394	0.376	0.449
Bi	2.284	2.226	2.368	2.290	2.376	2.289	2.249	2.240	2.278	2.318	2.386	2.354	2.322	2.312	2.236
S	7.146	7.077	7.132	7.179	7.191	7.161	7.108	7.154	7.087	7.167	7.089	7.035	7.099	7.082	7.091
Se	0.023	0.030	0.033	0.000	0.032	0.044	0.017	0.022	0.000	0.017	0.097	0.093	0.038	0.025	0.018
calcu	calculated empirical formulas are based on Pb+Bi+Sb = 5 apfu														









Fig. 9 Aikinite (dark gray) replaces nuffieldite (light gray) in quartz (black). BSE image by J. Sejkora.

Fig. 10 Aikinite (dark gray) replaces nuffieldite (light gray) in quartz (black). BSE image by J. Sejkora.



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	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15
Pb	34.06	35.32	33.84	34.49	34.11	33.90	34.00	34.39	33.62	34.40	34.02	34.70	34.99	34.48	34.86
Cu	10.61	10.37	10.49	10.49	10.26	10.29	10.21	10.44	10.00	10.53	10.44	9.56	9.95	10.45	10.33
Sb	0.46	0.56	0.57	0.25	0.18	0.25	0.32	0.23	0.56	0.54	0.38	1.05	1.39	0.47	0.63
Bi	37.06	37.19	37.20	36.63	37.88	37.89	37.94	36.84	39.01	36.89	37.50	36.76	35.88	37.48	37.45
S	17.25	17.29	17.42	16.87	17.17	17.19	17.09	17.00	17.70	17.15	17.34	17.29	17.22	17.20	17.42
Se	0.09	0.14	0.00	0.24	0.11	0.08	0.14	0.09	0.00	0.00	0.10	0.12	0.10	0.00	0.00
total	99.53	100.87	99.52	98.97	99.71	99.60	99.70	98.99	100.89	99.51	99.78	99.48	99.53	100.08	100.69
Pb	3.792	3.903	3.767	3.881	3.809	3.781	3.788	3.868	3.697	3.830	3.787	3.901	3.906	3.819	3.848
Cu	3.852	3.737	3.807	3.849	3.736	3.743	3.709	3.828	3.586	3.822	3.790	3.504	3.622	3.774	3.718
Sb	0.087	0.105	0.108	0.048	0.034	0.047	0.061	0.044	0.105	0.102	0.072	0.201	0.264	0.089	0.118
Bi	4.091	4.075	4.105	4.087	4.194	4.191	4.191	4.108	4.254	4.072	4.139	4.097	3.972	4.115	4.099
S	12.410	12.346	12.529	12.267	12.388	12.390	12.303	12.354	12.578	12.337	12.474	12.558	12.423	12.308	12.425
Se	0.026	0.041	0.000	0.071	0.032	0.023	0.041	0.027	0.000	0.000	0.029	0.035	0.029	0.000	0.000
n _{aik}	96	95	95	97	94	94	94	96	91	96	95	93	94	95	95
calcu	calculated empirical formulas are based on $(Cu+Pb)/2+(Sb+Bi) = 8 apfu$														

of nuffieldite from Slovinky given in Table 1 agrees well with the X-ray pattern calculated from the single-crystal data of nuffieldite from Les Houches in France (Moëlo et al. 1997); some observed differences in intensities of the individual diffraction maxima are caused probably by preferred orientation of the sample and small amount of sample available for experiment. The refined unit-cell parameters of studied nuffieldite are compared in Table 2 with published data.

The chemical composition of 11 different specimens of nuffieldite was quantitatively analysed and representative analyses as well as calculated empirical formulas are given in Table 3 (all 145 analyses are available in supplementary file). Moëlo (1989); Maurel, Moëlo (1990) and Moëlo et al. (1997) pointed that Sb content is essential for stability of nuffieldite in nature and excess of Cu is typical. Based on this they redefined empirical formula of nuffieldite as Cu1++- $Pb_2Bi_2(Pb_xSb_yBi_{1-x-y})S_7$, with x close to 0.37 and y ranging between 0.19 and 0.55. Aikinite-type Cu+Pb=(Bi, Sb)+vacancy substitution is characteristic for nuffieldite (Maurel, Moëlo 1990; Pršek et al. 2006; Pršek 2008). The x value in studied samples from Slovinky-Došťavná varies between 0.18 to 0.44 (with average of 0.30) and y value is ranging from 0.28 to 0.45 (with average of 0.40), which is in line with published data on chemical composition of nuffieldite (e.g. Moëlo 1989; Harris 1993; Mozgova et al. 1994; Moëlo et al. 1997; Pršek et al. 2006; Izumino et al. 2014). There is obvious negative correlation between Sb vs. Bi contents (Fig. 7), but contents of Cu vs. Sb has clearly positive trend (Fig. 8). Minor amounts of Se (reaching up to 0.10 apfu) were also detected. The average (n=145 analyses) empirical formula of nuffieldite from Slovinky-Došťavná based on Pb+Bi+Sb = 5 apfu is corresponding to $Cu_{130}Pb_{200}$ $\mathsf{Bi}_{2.00}(\mathsf{Pb}_{0.30}\mathsf{Bi}_{0.30}\mathsf{Sb}_{0.40})_{1.00}(\mathsf{S}_{7.12}\mathsf{Se}_{0.03})_{7.15}$

Aikinite forms microscopic, anhedral to subhedral grains or aggregates reaching up to 300 µm in size (Fig. 6, 9, 10), which are replacing nuffieldite. It is mostly associated with chalcopyrite, galena and native bismuth. Representative WDS analyses of aikinite from Slovinky are shown in Table 4 (all 29 analyses are available in supplementary file). Besides of dominant contents of Cu, Pb, Bi and S minor amounts of Sb (up to 0.26 *apfu*) and also Se (up to 0.07 *apfu*) were observed. The calculated hypothetical percentage of the aikinite end member n_{aik} is ranging from 91 to 97 (with average of 95). The average (n=29) empirical formula of studied aikinite based on (Cu+Pb)/2+(Sb+ Bi) = 8 *apfu* is Pb_{3.83}Cu_{3.76}(Bi_{4.12}Sb_{0.09})_{4.21}(S_{12.40}Se_{0.03})_{12.03}.

Conclusions

A new occurrence of rare Cu, Pb sulphosalt, nuffieldite was identified at the siderite-type hydrothermal carbonate-quartz vein in Došťavná near Slovinky. It is associated with chalcopyrite, pyrite, aikinite, galena, native bismuth, tourmaline and chlorite.

Acknowledgements

The authors wish to thank to Pavel Škácha for the microphotography of studied samples and to Zdeněk Dolníček for his kind help with the analytical work. This study was financially supported by the Ministry of Culture of the Czech Republic (long-term project DKRVO 2019-2023/1. I.c; National Museum, 00023272) and by VEGA project (2/0028/20).

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